User's guide to meteoland (ver. 0.7.1)

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1 Introduction to the package

1.1 Purpose

Reliable meteorological data are a basic requirement for hydrological and ecological studies at the landscape scale. Given the large spatial variation of meteorology over complex terrains, meteorological records from a single weather station are often not representative of entire landscapes. Studies made on multiple sites over a landscape require different meteorological series for each site; and other studies may require meteorological data series for all grid cells of a landscape, in a continuous way. In these cases, spatial correlation between the meteorology series of different sites or cells must be taken into account. For example, the sequence of days with rain of contiguous cells will normally be the same or very similar, even if precipitation amounts may differ. Finally, studies addressing the impacts of climate change on forests and landscapes require downscaling coarse-scale predictions of global or regional climate models to the landscape scale. When downscaling predictions for several locations in a landscape, spatial correlation of predictions is also important.

With the aim to assist research of climatic impacts on forests, the R package `meteoland` provides utilities to estimate daily weather variables at any position over complex terrains:

1. Spatial interpolation of daily weather records from meteorological stations.
2. Statistical correction of meteorological data series (e.g. from climate models).

Spatial interpolation is required when meteorology for the area and period of interest cannot be obtained from local sensors. The nearest weather station may not have data for the period of interest or it may be located too far away to be representative of the target area. Correcting the biases of a meteorological data series containing biases using a more accurate meteorological series is necessary when the more accurate series does not cover the period of interest and the less accurate series does. The less accurate series may be at coarser scale, as with climate model predictions or climate reanalysis data. In this case one can speak of statistical correction adn downscaling. However, one may also correct the predictions of climate models using reanalysis data estimated at the same spatial resolution.

1.2 Data structures

Package `meteoland` assists in the estimation of the following variables over lanscapes (units in parentheses):

- `DOY`: Day of the year ([1-365]).

- `MeanTemperature`: Mean daily temperature (in degrees Celsius).
- `MinTemperature`: Minimum daily temperature (in degrees Celsius).
- `MaxTemperature`: Maximum daily temperature (in degrees Celsius).
- `Precipitation`: Daily precipitation (in mm of water).
- `MeanRelativeHumidity`: Mean daily relative humidity (in percent).
- `MinRelativeHumidity`: Minimum daily relative humidity (in percent).
- `MaxRelativeHumidity`: Maximum daily relative humidity (in percent).
- `Radiation`: Incoming radiation (in MJ/m2).
- `WindSpeed`: Wind speed (in m/s).
- `WindDirection`: Wind direction (in degrees from North).
- `PET`: Potential evapo-transpiration (in mm of water).

The package deals with three kinds of spatial structures: individual points, a set of pixels from a spatial grid and full (i.e. complete) grids. The package includes six S4 spatial classes, which are defined as children of classes in package `sp`.

1.2.1 Topography

Three classes are defined to represent the variation of topographic features (i.e., elevation, slope and aspect) over space:

- `SpatialPointsTopography` extends `SpatialPointsDataFrame` and represents the topographic features of a set of points in a landscape.

  ```
  Class "SpatialPointsTopography" [package "meteoland"]

  Slots:

  Name:         data  coords.nrs      coords        bbox proj4string
  Class:  data.frame     numeric      matrix      matrix         CRS

  Extends:
  Class "SpatialPointsDataFrame", directly
  Class "SpatialPoints", by class "SpatialPointsDataFrame", distance 2
  Class "Spatial", by class "SpatialPointsDataFrame", distance 3
  ```

- `SpatialGridTopography` extends `SpatialGridDataFrame` and represents the continuous variation of topographic features over a full spatial grid.

  ```
  Class "SpatialGridTopography" [package "meteoland"]

  Slots:

  Name:          data         grid         bbox  proj4string
  Class:   data.frame GridTopology       matrix          CRS

  Extends:
  Class "SpatialGridDataFrame", directly
  Class "SpatialGrid", by class "SpatialGridDataFrame", distance 2
  Class "Spatial", by class "SpatialGridDataFrame", distance 3
  ```

- `SpatialPixelsTopography` extends `SpatialPixelsDataFrame` and represents the continuous variation of topographic features over a set if cells in a grid.

  ```
  Class "SpatialPixelsTopography" [package "meteoland"]

  Slots:

  Name:          data   coords.nrs         grid   grid.index
  Class:   data.frame      numeric GridTopology      integer

  Name:        coords         bbox  proj4string
  Class:       matrix       matrix          CRS

  Extends:
  Class "SpatialPixelsDataFrame", directly
  Class "SpatialPixels", by class "SpatialPixelsDataFrame", distance 2
  Class "SpatialPointsDataFrame", by class "SpatialPixelsDataFrame", distance 2
  Class "SpatialPoints", by class "SpatialPixelsDataFrame", distance 3
  Class "Spatial", by class "SpatialPixelsDataFrame", distance 4
  ```

Although the three classes have the same slots as their parent S4 classes, data frames in `SpatialPointsTopography`, `SpatialGridTopography` and `SpatialPixelsTopography` objects have only three variables: '**elevation**' (in meters), '**slope**' (in degrees) and '**aspect**' (in degrees from North).

1.2.2 Meteorology

Analogously to topography, three spatial classes are used to represent the variation of daily meteorology over space:

- SpatialPointsMeteorology extends SpatialPoints and represents daily meteorology series for a set of points in a landscape.

  ```
  Class "SpatialPointsMeteorology" [package "meteoland"]

  Slots:

  Name:        dates        data      coords        bbox proj4string
  Class:        Date      vector      matrix      matrix         CRS

  Extends:
  Class "SpatialPoints", directly
  Class "Spatial", by class "SpatialPoints", distance 2
  ```

- SpatialGridMeteorology extends SpatialGrid and represents the continuous variation of daily meteorology across a grid of cells.

  ```
  Class "SpatialGridMeteorology" [package "meteoland"]

  Slots:

  Name:          dates         data         grid         bbox
  Class:          Date       vector GridTopology       matrix

  Name:    proj4string
  Class:           CRS

  Extends:
  Class "SpatialGrid", directly
  Class "Spatial", by class "SpatialGrid", distance 2
  ```

- SpatialPixelsMeteorology extends SpatialPixels and represents the variation of daily meteorology for a set of pixels (cells) of a spatial grid.

  ```
  Class "SpatialPixelsMeteorology" [package "meteoland"]

  Slots:

  Name:          dates         data         grid   grid.index
  Class:          Date       vector GridTopology      integer

  Name:         coords         bbox  proj4string
  Class:        matrix       matrix          CRS

  Extends:
  ```

```
Class "SpatialPixels", directly
Class "SpatialPoints", by class "SpatialPixels", distance 2
Class "Spatial", by class "SpatialPixels", distance 3
```

In addition to their corresponding inherited slots, `SpatialPointsMeteorology`, `SpatialGridMeteorology` and `SpatialPixelsMeteorology` have two additional slots: '`dates`' (a vector of days specifying a time period), and '`data`' (a vector of data frames with the meteorological data). Although the three classes have a '`data`' slot containing data frames, meteorological data is in different form in each class. In objects of `SpatialPointsMeteorology`, there is one data frame for each point where variables are in columns and dates are in rows. In objects of `SpatialGridMeteorology` and `SpatialPixelsMeteorology`, each data frame describes the meteorology over a complete grid, or a subset of cells, for a single day. In these cases, the data frame has grid cells in rows and variables in columns.

1.3 Reading and writing meteorological data

1.3.1 Point meteorology

Objects of class `SpatialPointsMeteorology` are stored in the disk using one data file for each of their spatial points. Files can be stored in **ascii** format or **rds** (compressed) format. Package `meteoland` provides four input/output functions for point meteorology:

- Function `readmeteorologypoint()` reads the meteorological data stored in one **ascii/rds** data file and returns a data frame.
- Function `writemeteorologypoint()` writes the meteorological data of a single point as an **ascii/rds** file in the file system.
- Function `readmeteorologypointfiles()` reads several **ascii/rds** files and returns an object of class `SpatialPointsMeteorology`.
- Functions `writemeteorologypointfiles()` writes several **ascii/rds** files in the disk, one per spatial point. Metadata (i.e. the spatial coordinates of each point and the corresponding file path) is stored in an additional file.

1.3.2 Grid meteorology

Objects of class `SpatialGridMeteorology` are stored in the disk using one **netCDF** file per day. In this case, the **netCDF** file also contains the date and spatial projection. The following functions are available for input/output of meteorology over a full grid or on a subset of grid cells:

- Functions `readmeteorologygrid()` and `readmeteorologypixels()` reads the meteorological data stored in one **netCDF** file and returns an object of class `SpatialGridDataframe` or `SpatialPixels-Dataframe`, respectively.
- Functions `writemeteorologygrid()` and `writemeteorologypixels()` write the meteorological data of the full grid or the subset of grid cells, for a single date in a **netCDF** file.
- Functions `readmeteorologygridfiles()` and `readmeteorologypix-elsfiles()` read several **netCDF** files and returns an object of class `SpatialGridMeteorology` or `SpatialPixelsMeteorology`, respectively.
- Functions `writemeteorologygridfiles()` and `writemeteorologyp-ixelsfiles()` write several **netCDF** files in the disk, one per date, of the full grid or the subset of grid cells. Metadata (i.e. the dates and their corresponding file path) is stored in an additional file.

As a special case, function `readmeteorologygridcells()` reads several **netCDF** files and returns an object of class `SpatialPointMeteorology` with the meteorological data of a set of specified grid cells.

1.4 Visualizing meteorological data

Although very simple, the package provides two kinds of functions to visualize the temporal and spatial variation of meteorology:

- Function `spplot()` has been redefined from package `sp` to draw maps of specific weather variables corresponding to specific dates. The function can be used on objects of classs `SpatialGridMeteorology` and `SpatialPixelsMeteorology`.
- Function `meteoplot()` allows the temporal series of specific variables on specific spatial points to be plotted. The function can read the data from the disk or from objects of class `SpatialPointsMeteorology`.

1.5 Reshaping and summarizing meteorological data

The package provides a few functions to extract/reshape meteorological data:

- Function `extractgridpoints()` extracts the meteorology of particular points in a grid. It accepts objects of `SpatialGridMeteorology` and `SpatialPixelsMeteorology` as input and it returns an object of `SpatialPointsMeteorology`.

- Function `extractpointdates()` extracts the meteorology of a set of dates from an object of class `SpatialPointsMeteorology`. It returns one data frame for each date, with points in rows and meteorology variables in columns.

Other functions are used to generate temporal summaries of meteorological data. These accept meteorology objects as input and return their corresponding spatial dataframe structure with the summary statistics in columns:

- Function `summarypoints()` summarizes the meteorology of spatial points. It accepts objects of `SpatialPointsMeteorology` as input and returns an object of `SpatialPointsDataFrame` with point summaries for the requested variable. Temporal summaries can be calculated for different periods and using different summarizing functions (e.g. mean, sum, minimum, maximum, etc.).

- Functions `summarygrid()` and `summarypixels()` summarize the meteorology of full grids or of subset of grid cells, respectively. They accept objects of `SpatialGridMeteorology` and `SpatialPixelsMeteorology`, respectively, as input and return an object of `SpatialGridDataFrame` and `SpatialPixelsDataFrame`, respectively, with temporal summaries for the requested variable over the range of dates indicated.

1.6 Meteorology estimation functions

1.6.1 Spatial interpolation

Package `meteoland` provides two functions for interpolating meteorological data (i.e., one for each data structure):

- Function `interpolationpoints()` interpolates weather for a set of locations given in `SpatialPointsTopography` and returns an object of class `SpatialPointsMeteorology`.

- Function `interpolationgrid()` interpolates weather for a whole grid specified in `SpatialGridTopography` and returns an object of class `SpatialGridMeteorology`.

Both functions require an object of class '`MeteorologyInterpolationData`', which contains the X-Y coordinates, the meteorological data and topography of a set of weather stations as well as weather interpolation parameters.

```
Class "MeteorologyInterpolationData" [package "meteoland"]
```

```
Slots:

Name:                     coords                  elevation
Class:                    matrix                    numeric

Name:                      slope                     aspect
Class:                   numeric                    numeric

Name:             MinTemperature             MaxTemperature
Class:                    matrix                     matrix

Name:     SmoothedPrecipitation              Precipitation
Class:                    matrix                     matrix

Name:  SmoothedTemperatureRange           RelativeHumidity
Class:                    matrix                     matrix

Name:                  Radiation                  WindSpeed
Class:                       ANY                        ANY

Name:              WindDirection                 WindFields
Class:                       ANY                        ANY

Name:                    WFIndex                   WFFactor
Class:                       ANY                        ANY

Name:                     params                      dates
Class:                      list                       Date

Name:                       bbox                proj4string
Class:                    matrix                        CRS

Extends:
Class "MeteorologyProcedureData", directly
Class "Spatial", by class "MeteorologyProcedureData", distance 2
```

When calling functions `interpolationpoints()` or `interpolationgrid()`, the user may require interpolation outputs to be written into the file system, instead of being returned in memory. If `interpolationpoints()` is called with `export = TRUE`, the function will write the data frame produced for each point into an **ascii** text file or a **rds** file. If `interpolationgrid()` is called with `export = TRUE`, the function will write an **netCDF** file for each day. Metadata files will also be written, so that results can later be loaded in memory.

Functions `interpolation.calibration()` and `interpolation.cv()` are included in `meteoland` to calibrate interpolation parameters and evaluate predictive performance of interpolation routines before using them.

1.6.2 Statistical correction

One function is available for statistical correction of meteorological data series (i.e., one function for each data structure). Function `correctionpoints()` performs statistical correction of weather data series on a set of locations and it returns an object of class `SpatialPointsMeteorology` containing corrected weather predictions. Statistical correction requires an object of class '`MeteorologyUncorrectedData`', which contains the X-Y coordinates and the coarse-scale meteorological data to be corrected, which includes a reference (historic) period and projected (e.g. future) period:

```
Class "MeteorologyUncorrectedData" [package "meteoland"]

Slots:

Name:           coords  reference_data projection_data
Class:          matrix             ANY             ANY

Name:           params           dates            bbox
Class:            list            Date          matrix

Name:      proj4string
Class:             CRS

Extends:
Class "MeteorologyProcedureData", directly
Class "Spatial", by class "MeteorologyProcedureData", distance 2
```

The reference (historical) period is compared with observed meteorological data of the same period, and the routine uses this information to correct the projected (e.g. future) period. Therefore, apart from the '`MeteorologyUncorrectedData`' object, the correction function requires accurate meteorological data (for a set of spatial points or a grid). Normally, these data will be the result of spatial interpolation.

As before, when calling functions `correctionpoints()`, the user may require the outputs to be written into the file system, instead of being returned in memory. If `correctionpoints()` is called with `export = TRUE`, the function will write the data frame produced for each point into an **ascii** text file or a **rds** file. Metadata files will also be written, so that results can later be loaded in memory.

Function `correctionpoints.errors()` was included in the package to evaluate the errors of the less accurate and more accurate series. Comparisons can be made before and after applying statistical corrections. In the latter case, cross-validation is also available.

2 Spatial interpolation of weather records

2.1 Overview

Ecological research studies conducted for historical periods can be perfomed using meteorological records obtained from surface weather stations of the area under study. The general procedure for interpolation is very similar to the one that underpins the U.S. DAYMET dataset (https://daymet.ornl.gov/). For any target point, minimum temperature, maximum temperature and precipitation are interpolated from weather records using truncated Gaussian filters, while accounting for the relationship between these variables and elevation (Thornton et al. 1997). Relative humidity can be either interpolated (in fact, dew-point temperature is the variable being interpolated) or predicted from temperature estimates, depending on whether it has been measured in weather stations or not. Potential (i.e. top-of-atmosphere) solar radiation is estimated taking into account latitude, seasonality, aspect and slope, following Granier & Ohmura (1968). Potential solar radiation is then corrected to account for atmosphere transmittance using the predictions of temperature range, relative humidity and precipitation (Thornton & Running 1999). Finally, the wind vector (wind direction and wind speed) is interpolated by using weather station records and static wind fields.

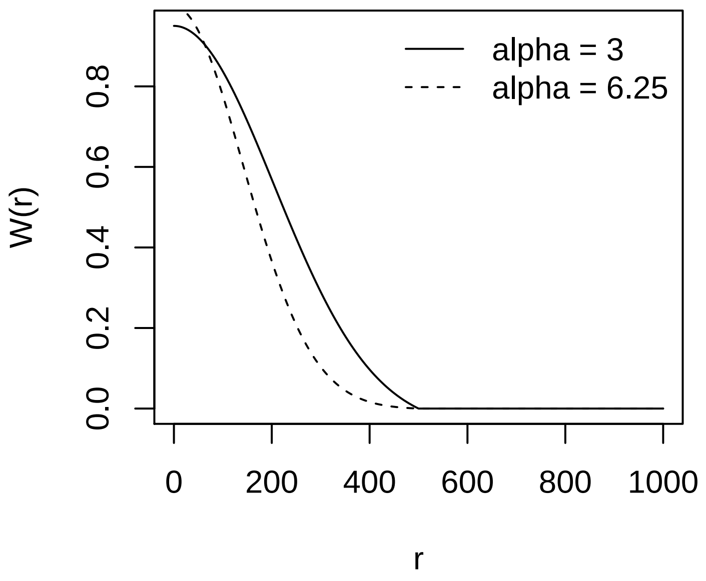
In the following subsections we detail the general algorithm used to obtain interpolation weights and the interpolation procedure for temperature, precipitation, relative humidity and wind. The estimation of potential and actual solar radiation is explained in next section.

2.2 Interpolation weights

Thornton et al. (1997) suggested interpolating meteorological data using a truncated Gaussian filter. Its form with respect to a central point p is:

$$W(r) = e^{-\alpha \cdot (r/R_p)^2} - e^{-\alpha} \tag{1}$$

if $r \leq R_p$ and $W(r) = 0$ otherwise. Here r is the radial distance from p, R_p is the truncation distance and α is the shape parameter. The spatial convolution of this filter with a set of weather station locations results, for each target point, in a vector of weights associated with observations. The following figure illustrates the Gaussian filter for $R_p = 500$ and either $\alpha = 3.0$ (continuous line) or $\alpha = 6.25$ (dashed line):



R_p is automatically adjusted so that it has lower values in data-rich regions and is increased in data-poor regions. The method, however, requires the user to specify N, the average number of observations to be included for each target point. R_p is then varied as a smooth function of the local density in such a way that this average is achieved over the spatial domain. Estimation of R_p is as follows:

1. A user-specified value is used to initialize R_p.
2. Interpolation weights W_i are calculated for all $i = (1, ..., n)$ stations, and the local station density is calculated as:

$$D_p = \frac{\sum_{i=1}^{n} (W_i/\hat{W})}{\pi \cdot R_p^2} \tag{2}$$

   where $\hat{W}$ is the average weight over the untruncated region of the kernel, calculated as:

$$\hat{W} = \left(\frac{1 - e^{-\alpha}}{\alpha} \right) - e^{-\alpha} \tag{3}$$

3. A new R_p value is calculated as a function of N and D_p, as:

$$R_p = \sqrt{\frac{N^*}{D_p \cdot \pi}} \tag{4}$$

where $N^* = 2N$ for the first $I - 1$ iterations, and $N^* = N$ for the final iteration.

4. The new R_p is substituted in step (2) and steps (2-4) are iterated a specified number of times I. The final R_p value is used to generate interpolation weights W_i.

Thornton et al. (1997) suggested to use this algorithm only once per point (and variable to be estimated), but since missing meteorological values can occur only in some days, we apply the algorithm for each target point and day. The interpolation method for a given set of observations is defined by four parameters R, I, N and α. Following Thornton et al. (1997), we set $R = 140000$ meters and $I = 3$ by default (see parameters '`initial_Rp`' and '`iterations`' given in function `defaultInterpolationParams()`). The other parameters (N and α) depend on the variable to be interpolated.

2.3 Temperature

Predictions for minimum temperature and maximum temperature are done in the same way, so we refer to a general variable T. We focus on the prediction of T_p, the temperature at a single target point p and for a single day, based on observations T_i and interpolation weights W_i for the $i = (1, ..., n)$ weather stations. Prediction of T_p requires a correction for the effects of elevation differences between observation points $z_1, ..., z_n$ and the prediction point z_p. Thornton et al. (1997) established the relationship between elevation and temperature using transformed variables (temporal or spatial moving window averages) for temperature and elevation, instead of the original variables, but we did not implement this feature here. A weighted least-squares regression is used to assess the relationship between temperature and elevation. Instead of regressing z_i on T_i, the independent variable is the difference in elevations associated with a pair of stations, and the dependent variable is the corresponding difference in temperatures. This gives a regression of the form:

$$(T_1 - T_2) = \beta_0 + \beta_1 \cdot (z_1 - z_2) \tag{5}$$

where subscripts 1 and 2 indicate the two stations of a pair and β_0 and β_1 are the regression coefficients. Regression is performed using all possible pairs of stations and the regression weight associated with each point is the product of the interpolation weights associated with the stations in a pair. The temperature for the target point, T_p is finally predicted as follows:

$$T_p = \frac{\sum_{i=1}^{n} W_i \cdot (T_i + \beta_0 + \beta_1 \cdot (z_p - z_i))}{\sum_{i=1}^{n} W_i} \tag{6}$$

where z_p is the elevation of the target point and z_i is the elevation of the weather station.

2.4 Relative humidity

Relative humidity is a parameter not always recorded in weather stations. When input station weather data does not include relative humidity, `meteoland` estimates it directly from minimum and maximum temperature (Thornton et al. 1997). Assuming that minimum daily air temperature $T_{min,p}$ at the target point is a good surrogate of dew-point temperature $T_{d,p}$ (i.e. $T_{d,p} = T_{min,p}$; note that this assumption may not be valid in arid climates), one can estimate actual vapor pressure e_p (in kPa) as:

$$e_p = 0.61078 \cdot e^{\left(\frac{17.269 \cdot T_{d,p}}{237.3 + T_{d,p}}\right)} \tag{7}$$

and saturated vapor pressure $e_{s,p}$ (in Pa) as:

$$e_{s,p} = 0.61078 \cdot e^{\left(\frac{17.269 \cdot T_{a,p}}{237.3 + T_{a,p}}\right)} \tag{8}$$

where $T_{a,p} = 0.606 \cdot T_{max,p} + 0.394 \cdot T_{min,p}$ is the average daily temperature. Finally, relative humidity RH_p (in percentage) is calculated as:

$$RH_p = 100 \cdot \frac{e_p}{e_{s,p}} \tag{9}$$

When relative humidity has been measured at weather stations, interpolation should be preferred to estimation from minimum and maximum temperature. However, because relative humidity depends on temperature, relative humidity RH_i of each weather station i has to be converted to dew-point temperature $T_{d,i}$ before interpolation (Tymstra et al. 2010). To obtain the dew-point temperature one first needs to calculate vapor pressure:

$$e_i = e_{s,i} \cdot (RH_i/100) \tag{10}$$

where $e_{s,i}$ is the saturated water vapor pressure of station i, calculated as indicated above. Then, dew-point temperature of station i is obtained from:

$$T_{d,i} = \frac{237.3 \cdot \ln(e_i/0.61078)}{17.269 - \ln(e_i/0.61078)} \tag{11}$$

Unlike temperature, interpolation of dew temperature is not corrected for elevation differences. The dew-point temperature for the target point, $T_{d,p}$ is predicted as:

$$T_{d,p} = \frac{\sum_{i=1}^{n} W_i \cdot T_{d,i}}{\sum_{i=1}^{n} W_i} \tag{12}$$

From the interpolated dew-point temperature one can obtain actual vapour pressure e_p and, together with saturated vapour pressure at point p, one calculates relative humidity as indicated above. If saturated vapour pressure is referred to average temperature $T_{a,p}$, then relative humidity is average relative humidity $RH_{a,p}$. If, instead, one refers saturated vapour pressure to minimum and maximum daily temperatures one obtains, respectively, maximum and minimum relative humidity values ($RH_{max,p}$, $RH_{min,p}$). After their estimation, the routine checks that the predicted maximum and minimum relative humidity values stay within the physical limits 0% and 100%. Although interpolation of dew-point temperature does not account for elevation differences, interpolated values of relative humidity will normally exhibit a pattern following elevation differences because temperature is involved in the calculation of relative humidity.

2.5 Precipitation

Predictions of precipitation are complicated by the need to predict both daily occurrence and, conditioned on this, daily precipitation amount. Thornton et al. (1997) define a binomial predictor of spatial precipitation occurrence as a function of the weighted occurrence at surrounding stations. The precipitation occurrence probability POP_p is:

$$POP_p = \frac{\sum_{i=1}^{n} W_{o,i} \cdot PO_i}{\sum_{i=1}^{n} W_{o,i}} \tag{13}$$

where PO_i is the binomial precipitation occurrence in station i (i.e., $PO_i = 0$ if $P_i = 0$ and $PO_i = 1$ if $P_i > 0$) and $W_{o,i}$ is the interpolation weight for precipitation occurrence. Once POP_p is calculated, then precipitation occurs if POP_p is smaller than a critical value (i.e. $PO_p = 1$ if $POP_p < POP_{crit}$ and $PO_p = 0$ otherwise).

Conditional on precipitation occurrence we calculate the prediction of daily total precipitation, P_p. Like with temperature, Thornton et al. (1997) established the relationship between elevation and precipitation using transformed variables (temporal or spatial moving window averages) for precipitation and elevation. Following their results, we transform precipitation values using a temporal window with side of 5 days. Weighted least-squares, where the weight associated with each point is the product of the interpolation weights associated with the stations in a pair, is used to account for elevation effects on precipitation. Unlike Thornton et al. (1997), who use the same set of interpolation weights (i.e. $W_{o,i}$) for precipitation occurrence and regression, we use a second set of interpolation weights $W_{r,i}$ for the calculation of regression weights. The dependent variable in the regression function is defined as the normalized difference of the precipitation observations P_i for any given pair of stations:

$$\left(\frac{P_1 - P_2}{P_1 + P_2}\right) = \beta_0 + \beta_1 \cdot (z_1 - z_2) \tag{14}$$

To obtain the predicted daily total P_p we use the following equation:

$$P_p = \frac{\sum_{i=1}^{n} W_{o,i} \cdot P_i \cdot PO_i \cdot \left(\frac{1+f}{1-f}\right)}{\sum_{i=1}^{n} W_{o,i} \cdot PO_i} \tag{15}$$

where $f = \beta_0 + \beta_1 \cdot (z_p - z_i)$. Note the usage of interpolation weight $W_{o,i}$ (and not $W_{r,i}$). The form of prediction requires that $|f| < 1$. A parameter f_{max} (with default $f_{max} = 0.95$) is introduced to force $|f| = f_{max}$ whenever $|f| > f_{max}$.

2.6 Wind

Interpolation of wind characteristics depends on the amount of information available:

- Interpolation of wind speed only
- Interpolation of wind vectors (speed and direction)
- Interpolation of wind vectors using wind fields

The following subsections detail the calculations in each case.

2.6.1 Interpolation of wind speed

The predicted wind speed u_p for a target point p is the weighted average of station wind speed values $\{u_i\}$ $i = (1, ..., n)$ using the interpolation weights W_i determined from the truncated Gaussian filter:

$$u_p = \frac{\sum_{i=1}^{n} W_i \cdot u_i}{\sum_{i=1}^{n} W_i} \tag{16}$$

2.6.2 Interpolation of wind vectors

Interpolation of wind vectors for a target point p is as follows. Let $\mathbf{v}_i$ be the wind vector in weather station i. $\mathbf{v}_i$ is initially expressed using polar coordinates. Indeed, we have u_i and θ_i, the wind speed and wind direction, respectively. If we express $\mathbf{v}_i$ in cartesian coordinates we have:

$$x_i = u_i \cdot \sin(\theta_i) \quad y_i = u_i \cdot \cos(\theta_i) \tag{17}$$

The predicted wind vector $\mathbf{v}_p$ is the weighted average of the wind vectors $\{\mathbf{v}_i\}$ $i = (1, ..., n)$ predicted for point p using the interpolation weights W_i determined from the truncated Gaussian filter:

$$x_p = \frac{\sum_{i=1}^{n} W_i \cdot x_i}{\sum_{i=1}^{n} W_i} \quad y_p = \frac{\sum_{i=1}^{n} W_i \cdot y_i}{\sum_{i=1}^{n} W_i} \tag{18}$$

The polar coordinates of the predicted wind vector $\mathbf{v}_p$ are:

$$u_p = \sqrt{x_p^2 + y_p^2} \quad \theta_p = \tan^{-1}(x_p/y_p) \tag{19}$$

2.6.3 Interpolation of wind vectors using wind fields

More precise wind interpolation of wind vectors requires a set of static wind fields covering the landscape of interest. Each of these wind fields has been calculated assuming a domain-level combination of wind speed and wind direction. The set of domain-level combinations should cover all possible winds in the landscape under study. For example, one could decide to include the combinations of eight different wind directions (i.e., N, NE, E, SE, ...) and three wind speed classes. The wind estimation of a given target point depends on both the wind observations at weather stations and these static wind fields.

In a given day (and before processing target points) we begin by identifying, for each weather station $i = (1, ..., n)$, the wind field m_i corresponding to a minimum difference between the observed wind vector $\mathbf{v}_i$ and the wind vector of the station in the wind field (i.e., minimum distance between the corresponding cartesian coordinates). The set of wind fields $\{m_i\}$ $i = (1, ..., n)$ chosen for each weather station conform the information for wind interpolation in a given day.

Actual wind interpolation details for a target point p are as follows. We first draw for each $i = (1, ..., n)$ the wind vector $\mathbf{v}_{m_i,p}$ corresponding to the location of the target point p in wind fields m_i. Let $u_{m_i,p}$ and $\theta_{m_i,p}$ be the wind speed and wind direction of $\mathbf{v}_{m_i,p}$, respectively. The cartesian coordinates of $\mathbf{v}_{m_i,p}$ are:

$$x_{m_i,p} = u_{m_i,p} \cdot \sin(\theta_{m_i,p}) \quad y_{m_i,p} = u_{m_i,p} \cdot \cos(\theta_{m_i,p}) \tag{20}$$

The predicted wind vector $\mathbf{v}_p$ is the weighted average of the wind vectors $\{\mathbf{v}_{m_i,p}\}$ $i = (1, ..., n)$ predicted for point p using the interpolation weights W_i determined from the truncated Gaussian filter:

$$x_p = \frac{\sum_{i=1}^{n} W_i \cdot x_{m_i,p}}{\sum_{i=1}^{n} W_i} \quad y_p = \frac{\sum_{i=1}^{n} W_i \cdot y_{m_i,p}}{\sum_{i=1}^{n} W_i} \tag{21}$$

The polar coordinates of the predicted wind vector $\mathbf{v}_p$ are found as before.

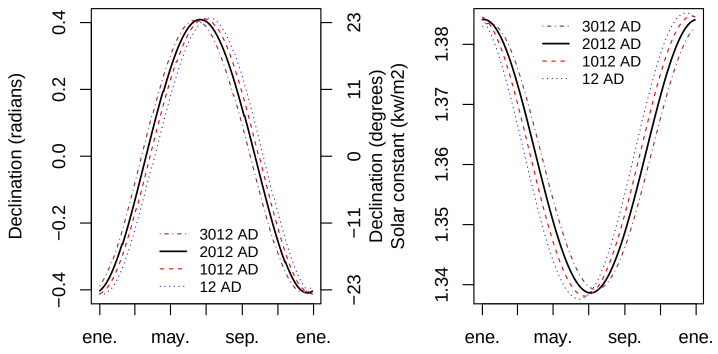
3 Estimation of solar radiation

Incident daily solar radiation is not interpolated, but estimated from topography and measurements of temperature, humidity and precipitation.

3.1 Solar declination and solar constant

The declination of the sun δ is the angle between the rays of the sun and the plan of the Earth's equator. Solar declination varies with years and seasons. However, the Earth's axial tilt changes slowly over thousands of years but it is nearly constant for shorter periods, so the change in solar declination during one year is nearly the same as during the next year. Solar constant (I_0) is normally given a nominal value of 1.361 kW·m^{-2} but in fact it also varies through the year and over years. Both can be calculated from Julian day (J), the number of days number of days since January 1, 4713 BCE at noon UTC. from Julian day. In `meteoland`, julian days, solar declination and solar constant are calculated using an adaptation of the code as in package `insol` (by J.G. Corripio), which is based on Danby (1988) and Reda & Andreas (2003).

The following figures show the variation of solar declination and the value of solar constant over a year (see functions `radiation_solarDeclination()` and `radiation_solarConstant()`):



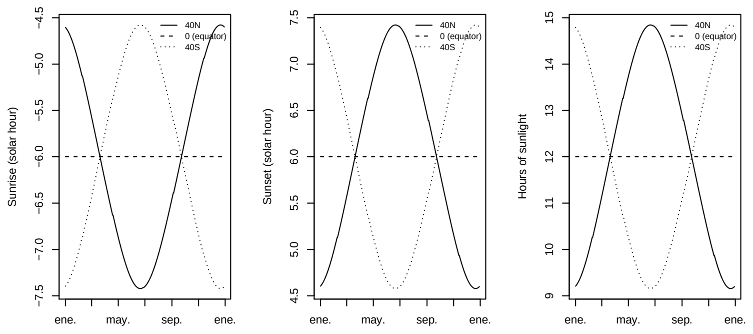
3.2 Day length

Calculation of sunrise and sunset on a horizontal surface is rather straightforward. The hour angles of sunrise and sunset (sr and ss, both in radians) for a horizontal surface of latitude ϕ on a day with declination δ (both expressed in radians) are:

$$sr = T_1 = \cos^{-1}(\max(\min(-\tan(\phi) \cdot \tan(\delta), 1), -1)) \quad (22)$$

$$ss = T_0 = -T_1 \quad (23)$$

Knowing that each hour corresponds to 15 degrees of rotation, hour angles can be transformed to solar hours. The following figures show the seasonal variation of sunrise and sunset hours, as well as day length, for horizontal surfaces in three latitudes (40North, equator and 40South) (see functions `radiation_sunRiseSet()` and `radiation_daylength()`):



For inclinated slopes, the calculation of day length is based on the concept of equivalent slopes, which are places on earth where the slope of earth's surface is equal to the slope of interest. The calculations start with the determination of the latitude L_1 of the equivalent slope:

$$\begin{aligned} L_1 &= \sin^{-1}\left(\cos(Z_x)\cdot\sin(\phi)+\sin(Z_x)\cdot\cos(\phi)\cdot\cos(A)\right) && (24)\\ D &= \cos(Z_x)\cdot\cos(\phi)-\sin(Z_x)\cdot\sin(\phi)\cdot\cos(A) && (25)\end{aligned}$$

where ϕ is the latitude, A is the azimuth of the slope (aspect) and Z_x is the zenith angle of the vector normal to the slope (equal to the slope angle). Then L_2 is defined depending on the value of D. If $D < 0$ then:

$$L_2 = \tan^{-1}\left(\frac{\sin(Z_x)\cdot\sin(A)}{D}\right)+\pi \qquad (26)$$

Otherwise, L_2 is calculated as:

$$L_2 = \tan^{-1}\left(\frac{\sin(Z_x)\cdot\sin(A)}{D}\right) \qquad (27)$$

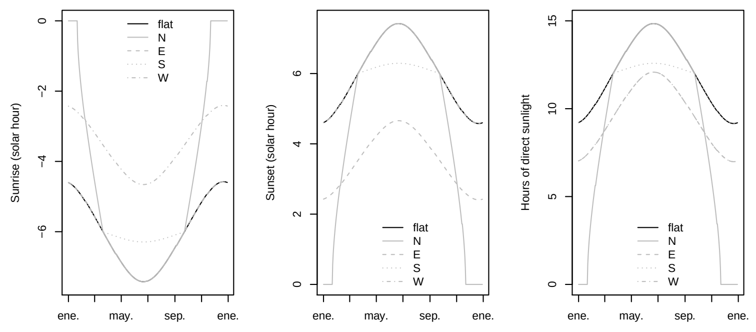
Once L_1 and L_2 are available, we can calculate solar hours on equivalent slopes:

$$\begin{aligned} T_7 &= \cos^{-1}\left(\max(\min(-\tan(L_1)\cdot\tan(\delta),1),-1)\right)-L2 && (28)\\ T_6 &= -\cos^{-1}\left(\max(\min(-\tan(L_1)\cdot\tan(\delta),1),-1)\right)-L2 && (29)\end{aligned}$$

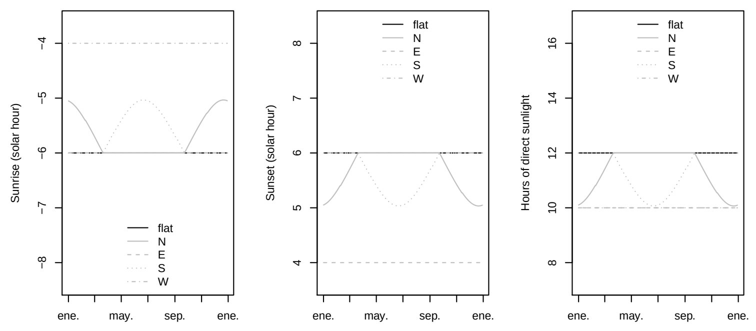
Being T_6 and T_7 the hour angle of sunrise and sunset on equivalent slopes, respectively. and the hour angles of sunrise (sr) and sunset (ss) on the slope (both in radians) are found comparing the hour angles on equivalent surfaces with the hour angles on the horizontal surface:

$$\begin{aligned} sr &= \max(T_0,T_6) && (30)\\ ss &= \min(T_1,T_7) && (31)\end{aligned}$$

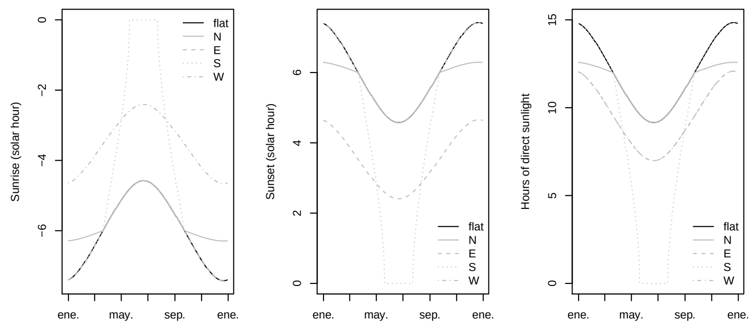
The following three figures show the seasonal variation of sunrise and sunset hours, as well as day length, for slopes of 30 inclination, facing to the four cardinal points. Curves for flat surfaces are shown for comparison. If the slopes are at latitude 40North:



whereas if they are at Equator (i.e. $\phi = 0$):



and if they are at latitude 40South:



3.3 Potential radiation

Potential solar radiation is the radiation that a surface on earth would receive if atmosphere was not present (i.e. without the effects of cloud reflection, scattering, ...). In `meteoland`, potential solar radiation is estimated from solar declination, latitude, aspect and slope according to Granier & Ohmura (1968). Daily potential radiation (R_{pot}, in MJ·m^{-2}) is calculated by integrating instantaneous potential radiation $R_{pot,s}$ (in kW·m^{-2}) over the day between sunrise (sr) and sunset (ss), using 10 min (i.e. 600 sec) intervals:

$$R_{pot} = \frac{1}{1000} \cdot \sum_{s=sr}^{ss} 600 \cdot R_{pot,s} \tag{32}$$

In turn, instantaneous potential solar radiation $R_{pot,s}$ is calculated using:

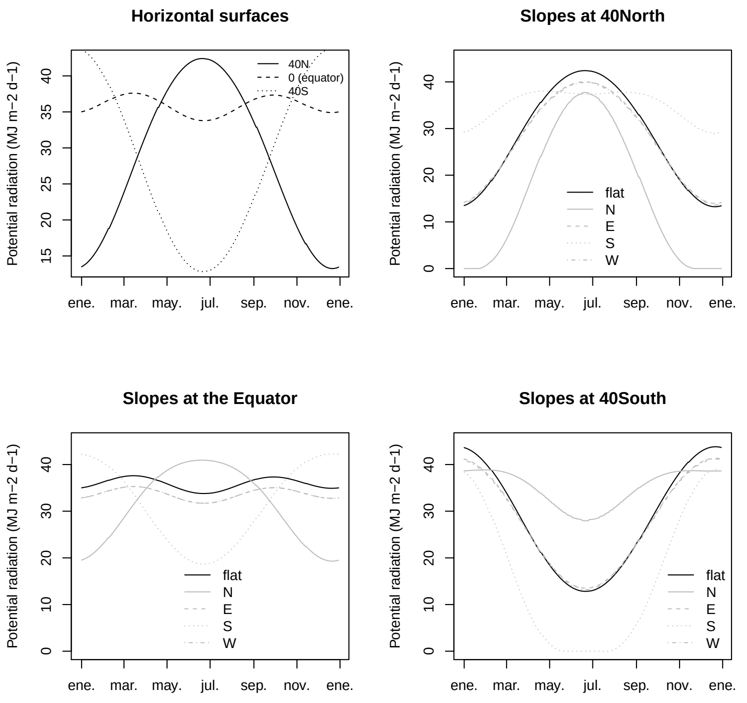
$$\begin{aligned} R_{pot,s} &= I_0 \cdot [(\sin\phi \cdot \cos H)(-\cos A \cdot \sin Z_x) - \sin H \cdot (\sin A \cdot \sin Z_x) \\ &\quad +[(\cos\phi \cdot \cos H) \cdot \cos Z_x] \cdot \cos\delta \\ &\quad +[\cos\phi \cdot (\cos A \cdot \sin Z_x) + \sin\phi \cdot \cos Z_x] \cdot \sin\delta] \end{aligned} \tag{33}$$

where I_0 is the solar constant, ϕ is the latitude, H is the hour angle measured from solar noon, positively towards the west, A is the azimuth of the slope (aspect), Z_x is the zenith angle of the vector normal to the slope (equal to the slope angle) and δ is the sun's declination. Note that in the case of a flat surface the previous equation reduces to:

$$R_{pot,s} = I_0 \cdot [\cos\phi \cdot \cos H \cdot \cos\delta + \sin\phi \cdot \sin\delta] = I_0 \cdot \sin\beta \tag{34}$$

where β is called the solar elevation angle.

The following figures illustrate seasonal variation of potential solar radiation for the horizontal inclinated surfaces presented above (see function `radiation_potentialRadiation()`):



3.4 Incident solar radiation

Incident solar radiation is the amount of (direct) solar radiation reaching the surface after accounting for the atmosphere. Improving the method proposed in Thornton et al. (1997), Thornton & Running (1999) calculate incident daily total solar radiation R_g as:

$$R_g = R_{pot} \cdot T_{t,max} \cdot T_{f,max} \tag{35}$$

where $T_{t,max}$ is the maximum (cloud-free) daily total transmittance and $T_{f,max}$ is the proportion of $T_{t,max}$ realized on a given day (cloud correction). The maximum daily total transmittance $T_{t,max}$ is estimated as:

$$T_{t,max} = \left[\frac{\sum_{s=sr}^{ss} R_{pot,s} \cdot \tau^{(P_z/P_0) \cdot m_\theta}}{\sum_{s=sr}^{ss} R_{pot,s}} \right] + (\alpha_{e_p} \cdot e_p) \tag{36}$$

where $\tau = 0.87$ is the instantaneous transmittance at sea level, at nadir, for a dry atmosphere; e_p is the actual water vapor pressure (in kPa), estimated as explained before; $\alpha_{e_p} = -6.1 \cdot 10^{-2}\text{kPa}^{-1}$ is a parameter describing the effect of vapour pressure on $T_{t,max}$; $m_\theta = 1/\cos\theta$ is the optical air mass at solar zenith angle $\cos(\theta) = \sin\phi \cdot \sin\delta + \cos\phi \cdot \cos\delta \cdot \cos H$; and P_z/P_0 is the ratio between air pressure at elevation z_p and air pressure at the sea level, calculated as:

$$(P_z/P_0) = (1.0 - 2.2569 \cdot 10^{-5} \cdot z_p)^{5.2553} \tag{37}$$

In turn, $T_{f,max}$ was empirically related to $\Delta T = T_{\max} - T_{\min}$, the difference between maximum and minimum temperatures for the target point:

$$T_{f,max} = 1.0 - 0.9 \cdot e^{-B \cdot \Delta T^C} \tag{38}$$

being $C = 1.5$ and B calculated from:

$$B = b_0 + b_1 \cdot e^{-b_2 \cdot \hat{\Delta T}} \tag{39}$$

with $b_0 = 0.031$, $b_1 = 0.201$ and $b_2 = 0.185$. In this last equation, $\hat{\Delta T}$ is a 30-day moving average for the temperature range ΔT. For computational reasons, we do not estimate $\hat{\Delta T}$ from the 30-day moving window average of predicted ΔT values, but from the interpolation of pre-calculated $\hat{\Delta T}$ values in weather stations. On wet days (i.e. if $P_p > 0$) the estimation of $T_{f,max}$ is multiplied by a factor of 0.75 to account for clouds.

Although the calculation of incident solar radiation can be done independently of interpolation (see function `radiation_solarRadiation()`), it is automatically done in functions `interpolationpoints()` and `interpolationgrid()`.

3.5 Outgoing longwave radiation and net radiation

Potential and actual evapotranspiration calculations require estimating the energy actually absorbed by evaporating surfaces. Daily net radiation R_n (in $MJ \cdot m^{-2} \cdot day^{-1}$) is calculated using:

$$R_n = R_s \cdot (1 - \alpha) - R_{nl} \tag{40}$$

where R_s is the input solar radiation (in $MJ \cdot m^{-2} \cdot day^{-1}$), $\alpha = 0.08$ accounts for surface albedo, and R_{nl} is the net longwave radiation. Outgoing longwave radiation is the radiation emitted by earth. Following McMahon et al. (2013) to obtain R_{nl} one first calculates clear sky radiation R_{so} using:

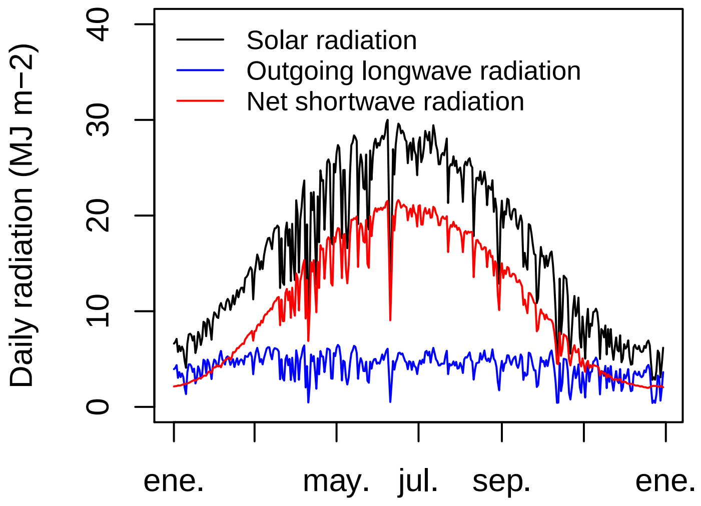
$$R_{so} = (0.75 + \cdot 0.00002 \cdot z) \cdot R_{pot} \tag{41}$$

where z is elevation and R_{pot} is potential radiation. R_{nl} is then calculated using:

$$R_{nl} = \sigma \cdot (0.34 - 0.14 \cdot \sqrt{e}) \cdot \frac{T^4_{\max} + T^4_{\min}}{2} \cdot (1.35 \cdot \min(\frac{R_s}{R_{so}}, 1.0) - 0.35) \tag{42}$$

where e is the actual vapor pressure (kPa), $T_{\max}$ and $T_{\min}$ are the maximum and minimum temperatures (in Kelvin) and $\sigma = 4.903 \cdot 10^{-9}$ MJ·K^{-4}·m^{-2} is the Stephan-Boltzmann constant.

The following figure shows an example of radiation balance for a whole year for a single site (see functions `radiation_outgoingLongwaveRadiation()` and `radiation_netRadiation()`):



3.6 Diurnal trends in diffuse and direct radiation

Ecological studies sometimes require radiation information at a subdaily scale. This is particularly true for modeling studies that need to calculate canopy photosynthesis. Although meteoland has been designed to assist studies requiring meteorological data at daily scale, a function called `radiation_directDiffuseDay()` is provided to divide daily radiation into instantaneous direct and diffuse radiation. Values of instantaneous direct and diffuse radiation (shortwave and photosynthetic active radiation) are calculated following Spitters et al. (1986). First, the ratio between daily diffuse and global radiation (R_d/R_g) is inferred from the ratio between daily potential and global radiation (R_g/R_{pot}):

$$R_d/R_g = 1 \qquad R_g/R_{pot} < 0.07 \tag{43}$$

$$R_d/R_g = 1 - 2.3 \cdot (R_g/R_{pot} - 0.7)^2 \qquad 0.07 \leq R_g/R_{pot} < 0.35 \tag{44}$$

$$R_d/R_g = 1.33 - 1.46 \cdot R_g/R_{pot} \qquad 0.35 \leq R_g/R_{pot} < 0.75 \tag{45}$$

$$R_d/R_g = 0.23 \qquad 0.75 \leq R_g/R_{pot} \tag{46}$$

In a clear day (e.g. not rainy) the ratio is modified to account for the circumsolar part of diffuse radiation:

$$R'_d/R_g = \frac{R_d/R_g}{1 + (1 - (R_d/R_g)^2) \cdot \cos^2(\pi/4 - \beta) \cdot \cos^3 \beta} \tag{47}$$

where β is the solar elevation angle. Otherwise $R'_d/R_g = R_d/R_g$. The daily diffuse shortwave radiation (R_d) is found by multiplying global radiation by the (modified) ratio:

$$R_d = R_g \cdot (R'_d/R_g) \tag{48}$$

The diurnal trend of the irradiance is derived from the daily global radiation and the daily course of potential (i.e. extra-terrestrial) radiation. If we assume that the atmospheric transmission is constant during the daylight period:

$$R_{g,s}/R_{pot,s} = R_g/R_{pot} \tag{49}$$

this leads to an estimation of the instantaneous global radiation (assuming compatible units):

$$R_{g,s} = R_g \cdot (R_{pot,s}/R_{pot}) \tag{50}$$

and the instantaneous diffuse and direct beam fluxes are estimated using:

$$R_{d,s} = R_d \cdot (R_{pot,s}/R_{pot}) \tag{51}$$

$$R_{b,s} = R_{g,s} - R_{d,s} \tag{52}$$

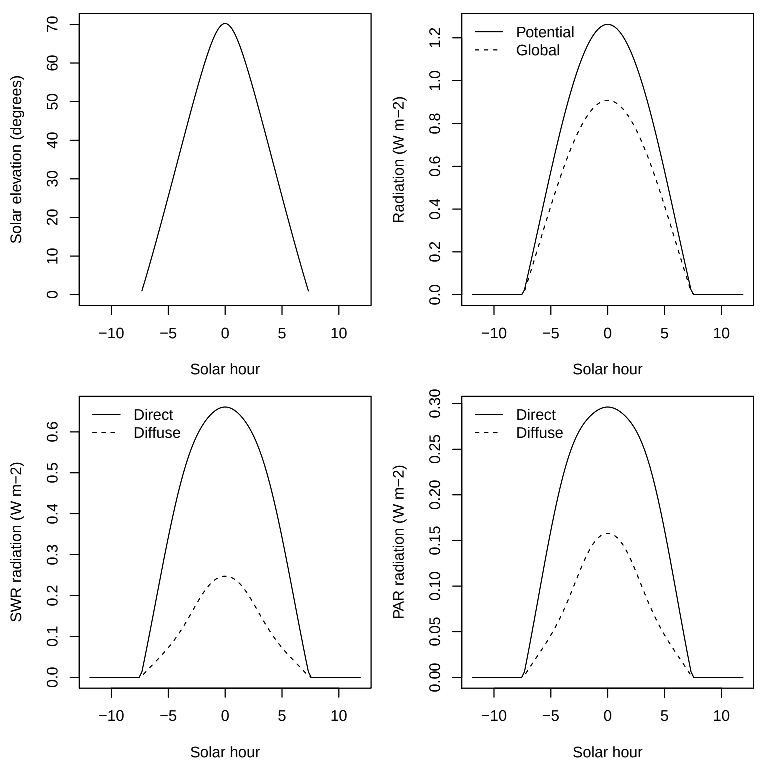
The whole procedure to calculate direct and diffuse radiation depends on the solar elevation angle, which changes through the day. Although R'_d/R_g is formulated as a ratio of daily values, the ratio needs to be calculated for every instant, as $R_{pot,s}$.

The procedure for photosynthetic active radiation (PAR) is similar. Daily PAR is assumed to be half of daily global radiation (i.e. $R_{PAR} = 0.5 \cdot R_g$. The scattered diffuse component of PAR is bigger than that of global radiation, and the ratio of diffuse over total PAR radiation is:

$$R_{PAR,d}/R_{PAR} = \left[1 + 0.3 \cdot (1 - (R_d/R_g)^2)\right] \cdot (R'_d/R_g) \tag{53}$$

The ratio $R_{PAR,d}/R_{PAR}$ is used to determine daily diffuse PAR and the calculation of instant rates are the same as for global radiation.

To illustrate the above calculations, we assume a target location in a flat terrain located at 42N latitude and 100 m.a.s.l, having 30 MJ$\cdot$m^{-2} of daily global radiation on the 2001/June/01, the hourly variation in diffuse and direct radiation would be:



4 Statistical correction of weather data

Statistical correction is necessary when meteorological data is available at a spatial scale that is too coarse for landscape-level analysis. This is usually the case when taking predictions from global or regional climate models. The general idea of correction to the landscape level is that a fine-scale meteorological series is to be compared to coarse-scale series for the a historical (reference) period. The result of this comparison can be used to correct coarse-scale meteorological series for other periods (normally future projections).

4.1 Correction methods

Users of `meteoland` can choose between three different types of corrections:

- **Unbiasing**: consists in subtracting, from the series to be corrected, the average difference between the two series for the reference period (Deque 2007). Let x_i be the value of the variable of the more accurate (e.g. local) series for a given day i and u_i the corresponding value for the less accurate series (e.g., climate model output). The bias, θ, is the average difference over all n days of the reference period:

$$\theta = \sum_{i}^{n}(u_i - x_i)/n \tag{54}$$

The bias calculated in the reference period is then subtracted from the value of u for any day of the period of interest.

- **Scaling**: A slope is calculated by regressing u on x through the origin (i.e. zero intercept) using data of the reference period. The slope can then be used as scaling factor to multiply the values of u for any day of the period of interest.
- **Empirical quantile mapping**: Due to its distributional properties, neither multiplicative or additive factors are appropriate for daily precipitation (Gudmundsson et al. 2012; Ruffault et al. 2014). In this case, it has been recommended to compare the empirical cumulative distribution function (CDF) of the two series for the reference period (Gudmundsson et al. 2012). The empirical CDFs of x and u for the reference period are approximated using tables of empirical percentiles, and this mapping is used to correct values of u for the period of interest:

  $$c_d = ecdf_x^{-1}(ecdf_u(u_d)) \tag{55}$$

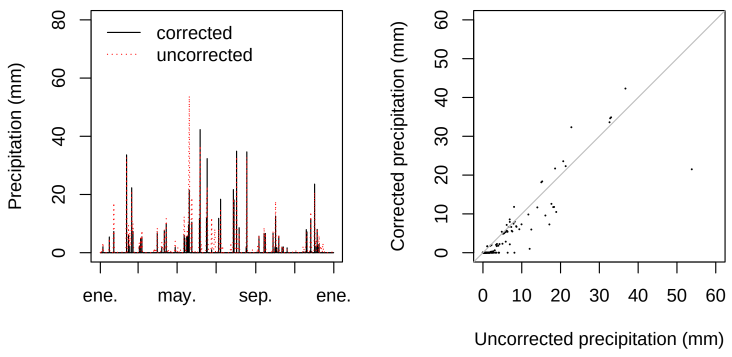
  where $ecdf_x$ and $ecdf_u$ are the empirical CDFs of x and u respectively. Values between percentiles are approximated using linear interpolation. `meteoland` performs quantile mapping by calling the functions implemented in package `qmap` (Gudmundsson et al. 2012).

For each target location to be processed, the correction routine first determines which is the nearest climate model cell and extracts its weather data series for the reference period and the period of interest. Then, the correction method chosen by the user for each variable is applied. Statistical corrections are done for each of the twelve months separately to account for seasonal variation of distributional differences (Ruffault et al 2014).

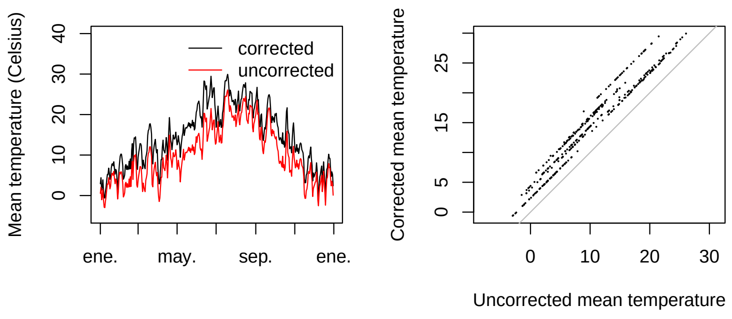
4.2 Default approaches by variable

Although users can choose their preferred correction method for each variable, meteoland has default approaches.

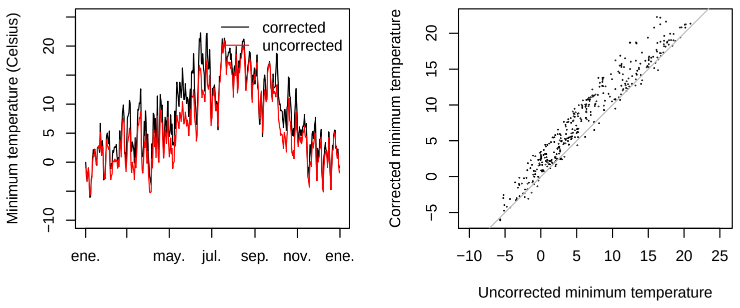
- **Precipitation**: By default, correction of precipitation is done using empirical quantile mapping. The following figures show the correction of RCM precipitation predictions for 2023 using interpolated data from 2002-2003 as observations for the reference period. Note that at least 15 years of observations (and not two!) would be needed for a correct estimation of monthly CDFs.

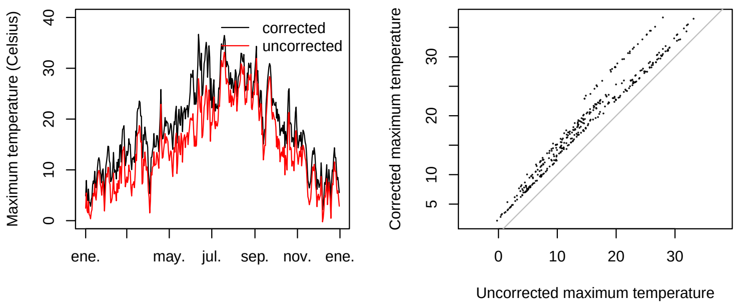


- **Mean temperature**: Unbiasing method is used by default to correct mean temperature. The following figures show the correction of RCM mean temperature predictions for 2023 using interpolated data from 2002-2003 as observations for the reference period:

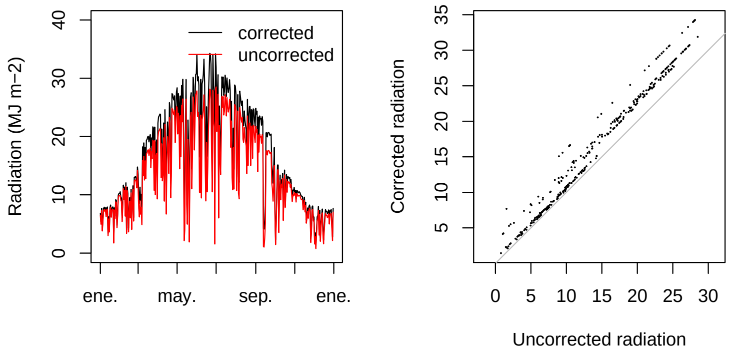


- **Minimum and maximum temperatures**: To correct minimum (respectively maximum) temperature values, by default scaling is applied to the difference between minimum (resp. maximum) temperature and mean temperature.





- **Radiation**: Radiation is by default corrected using the unbiasing procedure. The following figures show the correction of RCM radiation predictions for 2023 using interpolated data from 2002-2003 as observations for the reference period:



- **Relative humidity**: Mean relative humidity is first transformed to specific humidity, the unbiasing method is applied by default to this variable and the result is back transformed to mean, minimum and maximum relative humidity using the previously corrected series of mean, maximum and minimum temperature, respectively.
- **Wind speed**: By default, wind speed is corrected using the scaling method. Since historic wind data is often not available, however, if

wind speed data is missing the coarse-scale wind estimate is taken directly without correction.

5 Estimation of potential evapo-transpiration

Package `meteoland` allows calculating daily potential evapo-transpiration (PET) using Penman's formulation (Penman 1948; 1956) or Penman-Monteith formulation. PET is automatically calculated after meteorological data have been interpolated (i.e. within functions `interpolationpoints()` and `interpolationgrid()`) or downscaled (i.e. within functions `correction-points()` and `correctiongrid()`), but PET values can also be calculated for a single point using functions `penman()` or `penmanmonteith()`. For other formulations of PET, the reader is referred to package `'Evapotranspira-tion'`.

5.1 Penman formulation

Penman (1948) proposed an equation to calculate daily potential evaporation that combined an energy equation based on net incoming radiation with an aerodynamic approach. The Penman or Penman combination equation is:

$$E_{pot} = \frac{\Delta}{\Delta + \gamma} \cdot \frac{R_n}{\lambda} + \frac{\lambda}{\Delta + \lambda} \cdot E_a \tag{56}$$

where PET is the daily potential evaporation (in $mm \cdot day^{-1}$) from a saturated surface, R_n is the daily radiation to the evaporating surface (in $MJ \cdot m^{-2} \cdot day^{-1}$), Δ is the slope of the vapour pressure curve ($kPa \cdot {}^{\circ}C^{-1}$) at air temperature, γ is the psychrometric constant ($kPa \cdot {}^{\circ}C^{-1}$), and λ is the latent heat of vaporization (in $MJ \cdot kg^{-1}$). E_a (in $mm \cdot day^{-1}$) is a function of the average daily windspeed (u, in $m \cdot s^{-1}$), and vapour pressure deficit (D, in kPa):

$$E_a = f(u) \cdot D = f(u) \cdot (v_a^* - v_a) \tag{57}$$

where v_a^* is the saturation vapour pressure (kPa) and v_a the actual vapour pressure (kPa) and $f(u)$ is a function of wind speed, for which there are two alternatives (Penman 1948; 1956):

$$f(u) = 1.313 + 1.381 \cdot u \tag{58}$$

$$f(u) = 2.626 + 1.381 \cdot u \tag{59}$$

If wind speed is not available, an alternative formulation for E_{pot} is used as an approximation (Valiantzas 2006):

$$PET \simeq 0.047 \cdot R_s \cdot (T_a + 9.5)^{0.5} - 2.4 \cdot (\frac{R_s}{R_{pot}})^2 + 0.09 \cdot (T_a - 20) \cdot (1 - \frac{RH_{mean}}{100}) \tag{60}$$

where R_s is the incoming solar radiation (in $MJ \cdot m^{-2} \cdot day^{-1}$), T_a is the mean daily temperature (in °C), R_{pot} is the potential (i.e. extraterrestrial) solar radiation (in $MJ \cdot m^{-2} \cdot day^{-1}$) and RH_{mean} is the mean relative humidity (in percent).

5.2 Penman-Monteith formulation

The Penman-Monteith combination equation:

$$E_{pot} = \frac{1}{\lambda} \cdot \frac{\Delta \cdot R_n + D \cdot (\rho \cdot C_p / r_a)}{\Delta + \gamma \cdot (1 + r_c / r_a)} \tag{61}$$

where D is the vapour pressure deficit (in kPa), Δ is the slope of the saturated vapor pressure (in $Pa \cdot K^{-1}$), γ is the psychrometer constant (in $kPa \cdot K^{-1}$), λ is the latent heat vaporization of water (in $MJ \cdot kg^{-1}$) and C_p is the specific heat of air (in $MJ \cdot kg^{-1} \cdot K^{-1}$). r_c is the canopy resistance (in $s \cdot m^{-1}$). For simplicity, aerodynamic resistance (r_a) is currently set to $r_a = 208.0/u$ where u is the input wind speed.

6 Miscellaneous functions

6.1 Downloading data from AEMET

National meteorological agencies are increasingly adopting an open data philosophy and one of them is the Spanish meteorological agency (Agencia Española de Meteorología, AEMET). In `meteoland` we provide two functions that retrieve AEMET data using their OpenData API (API keys need to be obtained from AEMET):

- `downloadAEMEThistoricalstationlist()` : Gets the list of stations from which historical daily meteorological data is available.
- `downloadAEMEThistorical()` : Downloads historical daily meteorological data corresponding to a input set of stations and period.
- `downloadAEMETcurrentday()` : Downloads the last 24h of meteorological data corresponding to a input set of stations.

6.2 Extraction of climatic netCDF data

NetCDF is a standard data format for meteorological data. In particular, this format is used to store the predictions of global and regional climate models. Function `extractNetCDF()` parses a set of **NetCDFs** and extracts the daily meteorological data of a landscape boundary box for use within `meteoland`. The function first identifies which cells in NetCDF data should be extracted (according to the input boundary box), and the overall period.

For each cell to be processed, the function loops over all files (which can describe different variables and time periods) and extracts the corresponding data. The function transforms units to the units used in meteoland. If specific humidity and mean temperature are available, the function also calculates mean relative humidity.

6.3 Obtaining static wind fields

External software is necessary to calculate the set of wind fields for the study area under different domain-level average situations. For this we recommend using **WindNinja**, a computer program that calculates spatially varying wind fields for wildland fire applications. WindNinja allows simulating the spatial variation of wind for one instant in time. It was developed to be used by emergency responders within their typical operational constraints of fast simulation times (seconds), low CPU requirements (single processor laptops), and low technical expertise. WindNinja is typically run on domain sizes up to 50 kilometers by 50 kilometers and at resolutions of around 100 meters. The program is free and can be downloaded from www.firemodels.org.

The inputs for a basic run of WindNinja are an elevation data file for the study area, a domain-averaged input wind speed and direction and a specification of the dominant vegetation in the area. In order to obtain a set of pre-computed rasters of wind direction and speed, we suggest the following procedure:

- Export the elevation raster of the study area in one of the file formats accepted by WindNinja ('.asc', '.tif' or '.img'). In the case of a large study area (e.g. > 100 x 100 km) one should run WindNinja in subsets of the area and then integrate the results (e.g., Sanjuan et al. 2014).

- Run WindNinja, using the elevation of the study area, for all combinations of wind direction and wind speed class (for each wind speed class an mean class value has to be chosen). Several combinations of domain-level wind speed and wind direction can be specified for a single run, and the program can also be run in batch mode.

- Read raster files created by WindNinja (a wind speed file, a wind direction file) for each combination of domain-level wind speed and direction.

Function `readWindNinjaOutput()` can be used to conduct this last step. The function allows parsing all the ASCII raster files produced by WindNinja for combinations of wind direction (e.g., 0, 45, 90, 135, 180, 225, 270 and 315 degrees) and wind speed (e.g., 2, 7 and 10 m/s). The function returns a list with the following elements:

- The vector of domain-level wind directions corresponding to WindNinja Runs
- The vector of domain-level wind speed corresponding to WindNinja Runs
- An object `SpatialGridDataFrame` containing the wind directions (in degrees from North) for all WindNinja runs.
- An object `SpatialGridDataFrame` containing the wind speeds (in m/s) for all WindNinja runs.

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